SATELLITE DERIVED CLOUD-TRACK POLAR WINDS FROM 1981-2004

Richard Dworak⁺, Jeff Key^{*}, Dave Santek⁺, and Christopher Velden⁺

+ Cooperative Institute for Meteorological Satellite Studies, University of Wisconsin Madison, Wisconsin USA *NOAA/National Environmental Satellite, Data, and Information Service Madison, Wisconsin USA

ABSTRACT

Recent studies have shown that the Arctic climate has changed significantly over the past 20 years. Unfortunately, two important tools for studying recent climate change, the National Centers for Environmental Prediction (NCEP)/National Center for Atmospheric Research (NCAR) and the European Centre for Medium Range Weather Forecasts (ECMWF) reanalysis products, have been shown to have relatively large errors in the wind field over the Arctic where there is little or no radiosonde data available for assimilation. Can satellites be used to fill the spatial gaps? At least eight numerical weather prediction centers worldwide have demonstrated that satellite-derived polar winds have a positive impact on global weather forecasts. The impact on climate reanalyses should be similar. Therefore, a polar wind data set spanning more than 20 years is being generated using Advanced Very High Resolution Radiometer (AVHRR) data. Wind speed, direction, and height are estimated for the Arctic and Antarctic, poleward of approximately 65 degrees latitude, by tracking the movement of cloud features in the 11 μ m window channel. Validation of the satellite-derived winds is being performed with winds from radiosondes, the ECMWF reanalysis, and thermal winds from the TIROS Operational Vertical Sounder (TOVS). Results indicate that the AVHRR-derived winds have lower speed biases than ECMWF reanalysis when compared radiosondes. It is recommended that the historical AVHRR polar winds be assimilated in future versions of the reanalysis products.

1. INTRODUCTION

In the early 1960s, Tetsuya Fujita developed analysis techniques to use cloud pictures from the first TIROS polar orbiting satellite for estimating the velocity of tropospheric winds (Menzel, 2001). Throughout the 1970s and early 1980s, cloud motion winds were produced from geostationary satellite data using a combination of automated and manual techniques. During 1989 J. Turner and D.E. Warren from the British Antarctic Survey developed a manual and automatic based technique for deriving Infrared (IR) cloud track winds in the polar regions from sequences of AVHRR (Advanced Very High Resolution Radiometer) images from the polar orbiting TIROS-N series of satellites (Turner and Warren , 1989). In 1992, the U.S. National Oceanic and Atmospheric Administration (NOAA) began using an experimental automated winds software package developed at the University of Wisconsin Space Science and Engineering Center that made it possible to produce a full-disk wind set without manual intervention. Fully automated cloud-drift and water vapor motion vector production from the Geostationary Operational Environmental Satellites (GOES) became operational in 1996, and now wind vectors are routinely used in operational numerical models of the National Centers for Environmental Prediction (NCEP) (Nieman et al., 1997; Velden et al., 2005).

However, a major gap in the global observing system exists over the polar regions because no routine measurements of tropospheric winds are made over most of the Arctic Ocean and the Antarctic continent (Figure 1). Therefore, satellite-derived wind fields from polar orbiting satellites, such as the Moderate Resolution Imaging Spectroradiometer (MODIS) on-board Terra and Aqua satellites and AVHRR on-board NOAA satellites, are most valuable. Previous research has shown that the inclusion of MODIS winds has a positive impact on numerical weather forecasts (Key et al., 2003; Velden et al., 2005). For the purpose of climate reanalysis the AVHRR would be more useful than MODIS due to AVHRR having a longer record - 25 years (1981-2006) - compared to the much shorter 6-year (2000-2006) record of MODIS. With a lack of observational data over the polar regions, numerical weather prediction model forecasts and climate

reanalysis products will be less accurate over those regions. For example, Francis (2002) showed that both the NCEP and ECMWF climate reanalyses exhibited significant errors over the Arctic, being too strong by as much as 25-65% from the north and south directions when compared to radiosondes that were not assimilated into the reanalysis field (Francis 2002; Francis et al., 2005).

Recent studies have shown that the Arctic climate has changed significantly over the past 20 years. Examples of significant climate change over the polar regions include increased surface temperature over the Antarctic peninsula (Vaughan et al., 2001; Thompson and Solomon 2002), significant tropospheric warming over much of the Antarctic during the winter (Turner et al., 2006), increasing temperatures over most of the Arctic (Serreze et al. 2000; Rigor et. al., 2000; Cosimo 2003), decreased surface pressure (Walsh et al., 1996), and increased cyclonic activity in the Arctic (Zhang et al., 2004) and surrounding seas of the Antarctic (Fyfe 2003). As a result there has been decrease in sea ice concentration in the in the western Antarctic (Jacobs and Comiso 1997) and over much of the Arctic Ocean (Lindsay and Zhang 2005 and Rigor et al., 2002). Changes in these properties are a function of large-scale circulation patterns that affect surface-atmosphere interactions and feedback mechanisms. For example, an increase in cyclonic activity over the Arctic that has been observed since the 1980s has been a major influence in the thinning of sea ice and a decrease in ice concentration. It is therefore very important to improve any deficiencies in the reanalysis fields that are used for climate studies. Furthermore, improved wind reanalysis from satellitederived winds will most likely improve future short-term climate studies on trends in regional atmospheric wind and circulation patterns in the polar regions.



Figure 1: WMO stations across the Arctic and Antarctic. Only those stations that provide regular daily wind data are shown.

2. METHODOLOGY

Cloud tracking with AVHRR data is based on the established procedure used for the Geostationary Orbiting Environmental Satellites (GOES), which is essentially that described in Nieman et al. (1997) and Velden et. al. (1998). After remapping the orbital data to a polar stereographic projection, potential tracking features are identified. The lowest (coldest) brightness temperature in the infrared window band, generally indicating cloud, within a target box is isolated and local gradients are computed. Gradients that exceed a specified threshold are classified as targets for tracking.

Because AVHRR doesn't contain a water vapor channel, the wind vector heights are currently assigned by the infrared window method, which assumes that the mean of the lowest (coldest) brightness temperature values in the target sample is the temperature at the cloud top. This temperature is compared to a numerical forecast of the vertical temperature profile to determine the cloud height (Nieman et al., 1993). The method is reasonably accurate for opaque clouds, but inaccurate for semitransparent clouds (Nieman et al., 1993). In our research, the ECMWF ERA-40 reanalysis product with 2.5 degree spatial resolution and 13 vertical levels was used as the background field.

In order to calculate the wind vectors, a guess position of the target between images is determined and a search box of 32 by 32 pixels around the guess position is created. The target is then identified, if possible, in the search box of the subsequent image. The tracking method searches for the minimum of the sum of squared radiance difference between the target location and the region inside the search box (Key et al., 03). This is done for two image overlaps in three consecutive orbits, with the average of the two vectors taken as the initial calculated wind vector centered at the position of target in the middle image.

After wind vectors are determined and heights are assigned, the resulting data set is subjected to a rigorous post-processing, quality-control step. A 3-dimensional objective recursive filter is employed to re-evaluate the tropospheric level that best represents the motion vector being traced, to edit out vectors that are in obvious error, and to provide end users with vector quality information (Hayden and Purser 1995; Velden et al., 1998; Holmlund et al., 2001). An example of the final AVHRR graphical output product is given in Figure 2. When compared to the radiosonde network in Figure 1, it is obvious that the AVHRR winds can fill the gaps that are in the observational networks over the polar regions.



Figure 2: Example of AVHRR cloud-drift winds over the Arctic and Antarctic. The Arctic example (left) was taken from 1800 UTC on August 5, 1993. The Antarctic example was taken from 0600 UTC on April 25, 2001.

3. **RESULTS**

Statistical comparisons between AVHRR and ECMWF winds over the Arctic for random cases from 1992 through 2000 (Table 1) are given for three layers in terms of the wind speed root mean squared difference (RMS), the average speed and direction differences, and the mean wind speeds. The speed RMS is defined as the square root of the sum of the squared differences divided by the number of cases. The statistical comparisons between AVHRR and ECMWF indicate that the average speed differences change sign from negative (AVHRR being slower than ECMWF) in the lower and mid-levels to positive (AVHRR being faster than ECMWF) in the upper-levels, with the smallest speed differences occurring at mid-levels (400-700 hPa). The average direction differences, on the other hand, are negative (AVHRR counter-clockwise of ECMWF) for all layers and increasing with height. The speed RMS difference is about 44% of the mean ERA-40 wind speed. A statistical comparison of ERA-40 and AVHRR winds with radiosonde winds ("raobs") from the Leads Experiment (LeadEx, Beaufort Sea, 1992), showed similar results (Table 2). LeadEx radiosonde observations were not assimilated into the climate reanalysis and therefore provide an excellent measure on the accuracy of the two wind sets. The comparison to LeadEx indicated that the winds derived from AVHRR have overall lower speed bias and RMS difference than ECMWF, but a slightly higher direction bias than ECMWF. However, when compared by atmospheric layer, the AVHRR has a much lower direction bias than ECMWF at low levels. When compared to TOVS thermal derived winds (not shown), it was found that the correlation coefficient was in the range of 0.7 to 0.8 between AVHRR and TOVS winds. These results indicate that the AVHRR cloud-track winds have the potential to improve reanalysis products. However, more validation will be done in the near future to verify that AVHRR outperforms ERA-40.

Table 1. Statistics for AVHRR IR Winds Over the Arctic Compared to ECMWF.

Low-level(below 700 hPa)		All levels combined	
Speed rms	3.71 m/s	Speed rms ECMWF	5.99 m/s
Average speed difference	-0.92 m/s	Speed rms AVHRR	5.39 m/s
Average direction difference	-0.10 degrees	Speed bias ECMWF	1.64 m/s
Mean ECMWF speed	8.56 m/s	Speed bias AVHRR	0.71 m/s
Mean AVHRR speed	7.65 m/s	Direction bias ECMWF	2.02 degrees
		Direction bias AVHRR	3.55 degrees
Mid-level (400-700 hPa)		Mean ECMWF speed	8.81 m/s
Number of cases	72.542	Mean AVHRR speed	7.88 m/s
Speed rms	4.91 m/s	Mean raob speed	7.17 m/s
Average speed difference	-0.46 m/s		
Average direction difference	-0.70 degrees	Low-lovel (below 700 bBa)	
Mean ECMWF speed	11.01 m/s	Total collocations	70
Mean AVHRR speed	10.55 m/s	Speed rms ECMW/E	5.65 m/s
		Speed rms AV/HPP	5.00 m/s
		Speed his ECMWE	1.02 m/s
			0.27 m/s
High-level (above 400 hPa)		Direction bios ECMW/E	-0.37 m/s
Number of cases	9,993	Direction bias AV/HRR	-0.26 degrees
Speed rms	8.64 m/s	Moon ECMWE spood	-0.20 degrees
Average speed difference	1.71 m/s	Mean AV/HPP speed	6.21 m/s
Average direction difference	-1.53 degrees	Mean rach speed	6.68 m/s
Mean ECMWF speed	19.27 m/s	Mean raob speed	0.00 11/5
Mean AVHRR speed	21.07 m/s		
		Mid-level (400-700 hPa)	
		Total collocations	67
		Speed rms ECMWF	6.35 m/s
		Speed rms AVHRR	5.77 m/s
		Speed bias ECMWF	2.29 m/s
		Speed bias AVHRR	1.85 m/s
		Direction bias ECMWF	-6.45 degrees

 Table 2.
 Statistical Comparison of AVHRR and

 ECMWF to Raob Winds from LEADEX.
 Image: Comparison of AVHRR and

4. SUMMARY AND CONCLUSIONS

Previous research has shown that the reanalysis wind fields over the Arctic have significant errors that can have negative impacts on climate studies. MODIS satellite-derived winds have been proven to improve numerical weather forecasts, but satellite-derived winds have not yet been used to improve climate reanalysis products. With the climates over the Arctic and parts of the Antarctic changing significantly over the past twenty-five years, it is important to have very accurate wind fields for studying the mechanisms that gave rise to these changes.

Direction bias AVHRR

Mean ECMWF speed

Mean AVHRR speed

Mean raob speed

7.51 degrees

9.97 m/s

9.53 m/s

7.68 m/s

The AVHRR on board NOAA satellites provides a 25-year dataset that can be used to derive cloud-drift winds. Statistical analyses of the AVHRR winds have shown that overall the ECMWF ERA-40 wind speeds are faster than AVHRR, which is indication of the reanalysis being too strong as described by Francis (2002). Comparisons to LeadEx radiosonde data demonstrates that the ERA-40 has a significant higher speed bias than AVHRR winds. We therefore recommend that the AVHRR cloud-drift winds be used in future reanalysis products.

Acknowledgments. This project was supported by NOAA and NASA. We are grateful for the LeadEx data provided by Eli Hunter and Jennifer Francis of Rutgers University, and Ola Persson of the University of

Colorado. AVHRR data were obtained from NOAA's Comprehensive Large Array-data Stewardship System (CLASS) with the help of Aleksandar Jelenak. ERA-40 data were provided by ECMWF on the web. AVHRR data were calibrated and navigated with software provided by Dan Baldwin at the University of Colorado.

5. REFERENCES

- Comiso, J. C. (2003) Warming Trends in the Arctic From Clear Sky Satellite Observations. *Journal of Climate*, **16**, 21, pp 3498-3510.
- Francis, J. A. (2002) Validation of Reanalysis Upper-Level Winds in the Arctic with Independent Rawinsonde Data. *Geophysical Research Letters*, **29**, 9, pp 29-1-29-4.
- Francis, J. A., E. Hunter, and C.Z. Zou. (2005) Arctic Tropospheric Winds Derived From TOVS Satellite Retrievals. *Journal of Climate*, **18**, 13, pp 2270-2285.
- Fyfe, J. C. (2003) Extratropical Southern Hemisphere Cyclones: Harbingers of Climate Change? *Journal of Climate*, **16**, 17, pp 2802-2805.
- Hayden, C. M., and R. J. Purser. (1995) Recursive Filter Objective Analysis of Meteorological Fields: Applications to NESDIS Operational Processing. *Journal of Applied Meteorology*, **34**, 1, pp 3-15.
- Holmlund, K., C. S. Velden, and M. Rohn. (2001) Enhanced Automated Quality Control Applied to High-Density Satellite Winds. *Monthly Weather Review*, **129**, 3, pp 517-529.
- Jacobs, S. S., and J. C. Comiso. (1997) Climate Variability in the Amundsen and Bellinghausen Seas. *Journal of Climate*, **10**, 4, pp 697-709.
- Key, J. R., D. Santek, C. S. Velden, N. Bormann, J. Thepaut, L. P. Riishojgaard, Y. Zhu, and W. P. Menzel. (2003) Cloud-Drift and Water Vapor WInds in the Polar Regions From MODIS. *IEEE Transactions on Geoscience and Remote Sensing*, **41**, 2, pp 482-492.
- Lindsay, R.W., and J. Zhang. (2005) The Thinning of Arctic Sea Ice, 1988-2003: Have We Passed a Tipping Point? *Journal of Climate*, **18**, 22 pp 4879-4894.
- Menzel, W. P. (2001) Cloud Tracking with Satellite Imagery: From the Pioneering Work of Ted Fujita to the Present. *Bulletin of the American Meteorological Society*, **82**, 1, pp 33-47.
- Nieman, S. J., J. Schmetz, and W. P. Menzel. (1993) A Comparison of Several Techniques to Assign Heights to Cloud Tracers, *Journal of Applied Meteorology*, **32**, 9 pp 1559-1568.
- Nieman, S. J., W. P. Menzel, C. M. Hayden, D. Gray, S. T. Wanzong, C. S. Velden, and J. Daniels. (1997) Fully Automated Cloud-Drift Winds in NESDIS Operations. *Bulletin of the American Meteorological Society*, **78**, 6, pp 1121-1133.
- Rigor, I. G., J. M. Wallace, and R. L. Colony. (2002) Response of Sea Ice to the Arctic Oscilation. *Journal of Climate*, **15**, 18, pp 2648-2663.
- Rigor, I. G., R. L. Colony, and S. Martin. (2000) Variations in Surface Air Temperature Observations in the Arctic, 1979-97. *Journal of Climate*, **13**, 5, pp 896-914.
- Serreze, M. C., J. E. Walsh, F. S. Chapin III, T. Osterkamp , M. Dyurgerov, V. Romanovsky, W. C. Oechel, J. Morison, T. Zhang, and R.G. Barry. (2000) Observational Evidence of Recent Change in the Northern High-Latitude Environment. 1st ed. Vol. 46. The Netherlands: Kluwer Academic, pp 159-207.
- Thompson, D. W. J., and S. Solomon. (2002) Interpretation of Recent Southern Hemisphere Climate Change. *Science*, **296**, 5569, pp 895-899.
- Turner, J., and D. E. Warren. (1989) Cloud Track Winds in the Polar Regions From Sequences of AVHRR Images. *International Journal of Remote Sensing*, **10**, 4-5, pp 695-703.
- Turner, J., T. A. Lachlan-Cope., S. Colwell, G. J. Marshall, and W. M. Connolley. (2006) Significant Warming of the Antarctic Winter Troposphere. *Science*, **311**, 5769, pp 1914-1917.
- Vaughan, D. G., G. J. Marshall, W. M. Connolley, J. C. King, and R. Mulvaney. (2001) Devil in the Detail. *Science*, **293**, 5536, pp 1777-1779.
- Velden, C. S., C. M. Hayden, S. J. Nieman, W. P. Menzel, S. Wanzong, and J. S. Goerss. (1997) Upper-Tropospheric Winds Derived From Geostationary Satellite Water Vapor Observations. *Bulletin of the American Meteorological Society*, **78**, 2, pp 173-195.
- Velden, C., J. Daniels, D. Settner, D. Santek, J. Key, J. Dunion, K. Holmlund, G. Dengel, W. Bresky, and P. Menzel. Recent (2005) Innovations in Deriving Tropospheric Winds From Meteorological Satellites. Bulletin of the American Meteorological Society, 86, 2, pp 205-223.

Walsh, J. E., W. L. Chapman, and T. L. Shy. (1996) Recent Decrease of Sea Level Pressure in the Central Arctic. *Journal of Climate*, **9**, 2, pp 480-486.

Zhang, X., J. E. Walsh, J. Zhang, U. S. Bhatt, and M. Ikeda. (2004) Climatology and Interannual Variability of Arctic Cyclone Activity: 1948-2002. *Journal of Climate* **17**, 12, pp 2300-2317.